

Evaluation of Accuracy of the Geodetic Reference Systems for the Modelling of Normal Gravity Fields of Nigeria

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Abstract

The normal gravity fields of Nigeria have been modeled exploiting the Geodetic Reference System 1930 (GRS-30), Geodetic reference system 1967 (GRS-1967) and the world geodetic reference system 1984 (WGS-84). The determination of the normal gravity fields from the three Geodetic Reference Systems were carried out to evaluate the accuracy of one reference datum with respect to another. Descriptive statistics of the normal gravity field values modeled on regional and local scales showed a large difference ($16.11mGal$) between the 1930 reference earth model (GRS-30) and 1967 reference earth model (GRS-67). A large difference was also established between the 1930 and 1984 reference earth models. The large difference in the mean of normal gravity field is attributed to error inherent in the Potsdam value ($16mGal$). However, the small difference in normal gravity field values between the 1967 and 1984 reference systems is a pointer that the choice of either application in geophysical exploration or geodetic applications is of minor importance. The trends in discrepancies between the 1930 and 1967 reference earth models and 1930 and 1984 reference earth models are reflected in the standard deviation and standard error (S.E) of the normal gravity field values.

Keywords: Ellipsoid, Geodetic, Reference, Gravity, Normal, Statistics, Earth

1. Introduction

One of the stages in the analyses of gravity data is the conversion of gravity measurements to gravity anomalies. To determine gravity anomaly the theoretical (normal) gravity values are required. Arafin (2004) defined gravity anomaly at a point on the physical surface of the earth as the deviation of the earth's normal gravity value at the point of latitude from the observed gravity value at the same point. The normal gravity constitutes a theoretical value representing the acceleration of gravity that would be generated by a uniform ellipsoidal earth. If the figure of the earth is assumed to be ellipsoid of revolution, the normal gravity will vary with the latitude of the observation points. The ellipsoid of revolution can be accomplished by rotating an ellipse about its minor axis. Thus, at the poles, the earth is slightly flattened and bulges at the equator. Therefore, the ellipsoid of revolution is the geometrical figure used to approximate the shape of the earth. The undisturbed surface of the ocean (no wind, no current, no tides and averaged over time) perpendicular to the direction of the plumb line is called the geoid.

There are several factors that determine the value of normal gravity at a point on the ellipsoidal surface. They include the size and shape of the ellipsoid and a value computed from observational data which represents the

value of gravity at the equator. If normal gravity is computed for a relatively small area, for example a city; this constitutes local modelling of normal gravity. This type of modeling gives a local ellipsoidal surface (Fig.1a) and does not account for the curvature of the entire earth. The local ellipsoidal surface only closely approximates the size and shape of the earth in the small area. This idea represents that of a flat earth model. If a gravity survey is done for such small area (plane – table survey) exact location can be established relative to each other without accounting for the size and shape of the total earth. The difference in gravity value between the measured gravity and normal gravity values computed for such a relatively small area constitutes gravity anomaly. Determination of different types of gravity anomalies require that the observed data be reduced to an equipotential surface called the geoid.

The computation of normal gravity for the entire earth for the determination of the ellipsoid constitutes global modeling (Fig.1a). Where the local ellipsoidal surface is large, the modelling of the normal gravity field for example, country/region constitutes regional modelling. Osazuwa (1993) determined the normal gravity values for Nigeria and evaluated the accuracy of the geodetic reference systems (GRS-30, GRS-67 and GRS-80). This study considered an additional reference earth model, the WGS-84 (World geodetic system 1984) which represent the latest reference earth model for determination of normal gravity field. Therefore, the objective of this study is to compute the normal gravity field of Nigeria per $0.5^\circ \times 0.5^\circ$ change in latitude (regional modeling) and normal gravity field for a geological province (fig1b) in southeastern Nigeria (local modeling) where gravity data exists. Thus, where gravity observation exists, normal gravity field values could be handy for gravity anomalies determination. Gravity anomalies are useful for tectonic studies, structural relations of deep and adjacent geologic features (Okiwelu et al., 2010) and basin analyses. It is also useful for anomaly transformation (Saheel et al., 2010) and reconnaissance tool for hydrocarbon exploration. Since normal gravity field data are indispensable in modelling gravity anomalies the reference earth models for determining normal gravity fields should be evaluated to ascertain their accuracy. This is the main objective of this study. Normal gravity fields are computed from International gravity formula (IGF) [Geodetic Reference Systems (GRS)]. The popular reference earth models are the GRS-67 (Geodetic Reference System 1967) and WGS-84 (World Geodetic Reference System 1984). The 1930 formula (GRS-30) was previously used. Other earlier Schemes include WGS-72, WGS-66 and WGS-60. These geodetic Systems e.g. WGS84 are standard for coordinate frame for the earth and a standard spheroidal reference ellipsoid for raw data and gravitational equipotential surface that defines the nominal sea level. The geodetic systems are therefore very suitable for computing the equipotential surfaces. The WGS-84 reference ellipsoid is more amenable to geophysical applications (Fairhead et al., 2003, Kuhn et al., 2009) while the GRS-80 is usually used for geodetic applications. The earth's physical surface comprising the mountains, valleys, rivers and surface of the sea is highly irregular and not suitable as a computational surface. The geoid represents a smooth surface while the ellipsoid is a better smooth mathematical surface that best fits the shape of the earth. Both equipotential surfaces do not coincide. The separation between the two is referred to as geoid undulation or geoid heights.

2. Theory and Methodology

2.1 Theory

The gravity on the ellipsoid can be derived from the gravitational potential U:

$$U = \frac{GM}{r} + \frac{GMa^2}{r^3} J_2 \left[\frac{3}{2} \sin^2 \phi - \frac{1}{2} \right] - \frac{1}{2} \omega^2 r^2 \cos^2 \phi \quad (1)$$

Where r is the radius of the spheroid and varies with latitude, ϕ according to

$$r(\phi) = a(1 - f \sin^2 \phi) \quad (2)$$

a is the radius of the earth at equator while G , M , ω are the gravitational constant, mass of the earth and the angular speed of the earth rotation respectively. The second term in equation (1) is due to spheroidal shape of the earth. J_2 is a constant determined by the distribution of mass and the term in bracket is the second degree harmonic giving the spheroidal shape. The third term is the centrifugal potential. f in equation (2) is the flattening of the earth.

$f = \frac{a-c}{a} = \frac{1}{298.257223560}$ (Sandwell, 2002), where c is the radius of the earth at the pole. Thus, the gravity on

the ellipsoid is

$$g = -\nabla u \sqrt{\left(\frac{\partial u}{\partial r}\right)^2 + \left(\frac{\partial u}{r\partial\theta}\right)^2} \quad (3)$$

$$\approx \frac{\partial u}{\partial r} = \frac{GM}{r^2} - \frac{3GMa^2}{r^4} J_2 \left[\frac{3}{2} \sin^2 \phi - \frac{1}{2} \right] - \omega^2 r \cos^2 \phi$$

The value of gravity on the ellipsoid is the normal gravity γ given by

$$\gamma(\phi) = g_{eq} [1 + p \sin^2(\phi) + q \sin^2(2\phi)] \quad (4)$$

$$g_{eq} = 978.0327 \text{ gal (gravity at equator)}$$

Where

$$\rho = 0.0053024$$

$$q = -0.0000059$$

The ellipsoidal model can now be written for GRS30, GRS1967, GRS80 and WGS84. Thus,

$$\gamma_{(\phi)1930} = 978049.0 \left(1 + 0.0052884 \sin^2 \phi - 0.0000059 \sin^2 2\phi \right) mGal \quad (5)$$

$$\gamma_{(\phi)1967} = 978031.8 \left(1 + 0.0053024 \sin^2 \phi - 0.0000059 \sin^2 2\phi \right) mGal \quad (6)$$

$$\gamma_{(\phi)1980} = 978032.7 \left(1 + 0.0053024 \sin^2 \phi - 0.0000058 \sin^2 2\phi \right) mGal \quad (7)$$

$$\gamma_{(\phi)1984} = 978031.85 \left(1 + 0.00527889 \sin^2 \phi + 0.000023462 \sin^4 \phi \right) mGal \quad (8)$$

The expression for gravity anomaly can be written as

$$\Delta g = g_{obs} - \gamma_{\phi} \quad (9)$$

Where g_{obs} is observed gravity

2.2 Methodology

The ellipsoidal model of the earth listed in equations (5, 6, 7 and 8) are utilized in geodesy for computing the shape of the earth and applied in exploration geophysics for determination of gravity anomalies. At the 1967 meeting of the International Union of Geodesy and Geophysics (IUGG) held in Lucerne Switzerland, the ellipsoid (GRS-67) was recommended for use to make up for the accuracy and precision lacking in GRS-30 which was adopted at the general assembly in Stockholm in 1930. The GRS67 was eventually approved and adopted at the 1971 meeting of the IUGG held in Moscow and therefore replaced the earlier Geodetic reference systems.

In December, 1979 at Canberra the general assembly of the IUGG adopted the Geodetic Reference System 1980. The body recognized the fact that the Geodetic reference System 1967 no longer represent the size, shape and gravity field of the earth to an accuracy adequate for many geodetic, geophysical, astronomical and hydrographic applications (Moritz, 1980). The GRS-67 was then replaced by a new GRS-80 based on the theory of the geocentric equipotential ellipsoid defined by the following Conventional Constants (Moritz, 1980): equatorial radius of the earth (a) = 6378137m, geocentric gravitational constant of the earth (including atmosphere, $GM = 3986005 \times 10^8 \text{ m}^3 \text{ s}^{-2}$, dynamical form factor of the Earth, excluding the permanent tidal deformation:

$$J_2 = 108263 \times 10^{-8}, \text{ angular velocity of the Earth: } \omega = 7292115 \times 10^{-11} \text{ rad/s}$$

Due to improved data, increasing data coverage, new data types and improved techniques, the idea of new WGS emerged. Using the GRS-80 parameters in addition to available doppler, satellite laser ranging and very long Baseline interferometry (VLBI), the WGS-84 (World Geodetic System 1984) was conceptualized. Additional data from satellite radar altimetry and sophisticated technique such as least squares method (Collocation) presented a uniform combination solution from different types of measurements; all relative to the earth's gravity field; including geoid and gravity anomalies. This latest earth gravitational model is currently the reference System being used by the Global Positioning Systems. At present the WGS84 exploits the 1996 Earth's Gravitational model (EGM96) geoid which was revised in 2004. This geoid defines the nominal sea level surface by means of spherical harmonics series of degree 360.

The normal gravity fields for Nigeria ($lat.2^{\circ} - lat.15^{\circ}$) have been determined based on the GRS-30, GRS-67 and WGS-84 using QCTools Software. An example of local modeling of normal gravity field is demonstrated using geological province (Calabar Flank) [fig.1b] in the Southeastern Sector of Nigeria. To evaluate the accuracy of the computed normal gravity fields, the values were subjected to descriptive statistical analyses using MINITAB 14 software.

3. Results and Discussion

The results of the normal gravity fields computed from the GRS-30, GRS-67 and WGS-84 on a regional and local scale are presented in *Tables 1 and 2* respectively. It is important to look at the descriptive statistics (e.g. mean, standard deviation) [table3, table4, fig.2, fig.3, fig.4, fig.5, fig.6 and fig.7]. The local modeling, for example, gives a mean difference of $16.79mGal$ between the GRS-30 and GRS-67. This result is consistent with the changes in theoretical gravity values with a difference of about $17.00mGal$ at the equator. However, there is a minimal difference of $1.3603mGal$ between WGS-84 and GRS-67. The conversion from 1967 reference model to 1984 reference model (Osazuwa, 1993) gives a value of $-1.3603mGal$. The large difference in values between GRS-30/GRS-67 and GRS-30/GRS-84 is a manifestation of the error in the Potsdam value (Dryden, 1942; Jeffreys, 1949; Morelli, 1959 and Wollard, 1950). This error necessitated a change from 1930 reference model to 1967 reference model. There was a recognition that the absolute gravity values at Potsdam System (1930 model) base station were in error and that allowing for an improved knowledge of the shape of the earth required a corresponding change in formula (Milsom, 2003).

The mean values of normal gravity field for Nigeria (regional modeling) are $978167 \pm 20.4mGal$ and $978169 \pm 20.4mGal$; the values were computed from GRS-67 and WGS84 respectively. The closeness in values suggests that the choice of either approach is of minor importance for most applications in geophysical exploration and geodesy. The change from 1967 reference model to 1984 world geodetic system is progressing gradually since the actual changes implied in theoretical gravity are often smaller than the errors in absolute gravity of individual gravity stations and no changes in base station values are required (Milsom, 2003). Despite the change from GRS-67 to WGS-80 and now WGS-84, the 1967 International Gravity formula is still very popular because of its compatibility with the network of international gravity base stations known as IGSN71 [International Gravity Standardization Network(1971)]. While the International gravity formula (geodetic reference systems) serve as a means of computing the normal gravity at the reference ellipsoid based on the latitude of the point of consideration, the International standardization network serves as a reference datum by means of which absolute gravity values are extrapolated to another area through relative gravity measurements (Osazuwa, 2004). The changes in geodetic reference system in computing the normal gravity fields is due to changes in some geodetic parameters such as Newtonian gravitational Constant(G), Geocentric gravitational Constant(GM), mean equatorial gravity in the Zero-frequency tide system (g_{eq}) and equatorial radius of the reference ellipsoid(a). The WGS-84 was computed by considering the variation of atmospheric density. This is done by applying correction to measured values (Moritz, 1980). The reference ellipsoid is defined to enclose the whole mass of the earth including the atmosphere. The 1967 geodetic reference system computation is, however, based on the theory of the equipotential ellipsoid without an atmosphere.

The results of local modeling in *Table 4* from the three reference models show that the standard deviations ($2.54mGal$) are low. This suggests that the set of data were computed at close interval. This is approximately at 0.01° intervals. The computation of data point at $0.5^{\circ} \times 0.5^{\circ}$ change in latitude for the regional modeling increased the standard deviation to $106mGal$. The similarity in standard deviations of $2.54mGal$ and $106mGal$ for the local and regional modeling respectively for the three reference models is a pointer that the data sets were computed at equal intervals. This interpretation is compatible with the standard errors in *Tables 3 and 4*. The

standard errors ($0.158mGal$) for the three reference models are similar for the locally modeled normal gravity values and $20.4mGal$ for the regional modeled values.

The relationship between the three reference models is demonstrated in *Figs.5 and 6* (normal curve superimposed in histogram). The curve is skewed more to the left in the 1930/1967 and 1930/1984 reference earth models than the 1967/1984 reference earth models. These relationships are revealed also in the box plot. Medians for normal gravity fields for Nigeria are $978145mGal$ and $978146mGal$ for 1967 and 1984 geodetic reference systems respectively. This implies a minimal difference and close association between the two models. The median value for the 1930 model is however $978162mGal$ suggesting a large difference from the other models. This relationship is propagated to the locally modeled values *Figs. 6 and 7* which are not unconnected with the error in Potsdam value.

4. Conclusion

Evaluation of the normal gravity field determined from the geodetic reference systems (1930, 1967 and 1984) for Nigeria revealed a large difference between the 1930 and 1967 reference earth models and 1930 and 1984 reference earth models. The large difference in the means of normal gravity fields manifests as constant bias ($16mGal$). The large difference is consistent with the error in Potsdam value. The minor difference between the values computed from the GRS-67 and WGS-84 suggests that the choice of either approach is of minor importance for most applications in geophysical exploration and geodetic applications. The change in geodetic reference systems in computing normal gravity fields is due to changes in some geodetic parameters. The results of the normal gravity fields are consistent because the values are dependent only on latitude of the computation point. The standard deviation and standard errors are also consistent indicating that the data points were computed at equal interval.

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Table 1. Regional modeling of normal gravity fields from the geodetic reference systems

Station No.	Latitude (Deg.)	GRS-30 (mGal)	GRS-67 (mGal)	WGS-84 (mGal)
1	2.000	978055.272	978038.138	978039.502
2	2.500	978058.797	978041.673	978043.045
3	3.000	978063.104	978045.992	978047.368
4	3.500	978068.191	978051.092	978052.456
5	4.000	978074.057	978056.973	978058.345
6	4.500	978080.699	978063.633	978065.010
7	5.000	978088.116	978071.070	978072.434
8	5.500	978096.305	978079.281	978080.620
9	6.000	978105.264	978088.264	978089.612
10	6.500	978114.991	978098.017	978099.367
11	7.000	978125.482	978108.536	978109.894
12	7.500	978136.735	978119.818	978121.163
13	8.000	978148.745	978131.861	978133.185
14	8.500	978161.510	978144.659	978146.004
15	9.000	978175.025	978158.210	978159.556
16	9.500	978189.286	978172.510	978173.859
17	10.000	978204.289	978187.553	978188.915
18	10.500	978220.030	978203.336	978204.703
19	11.000	978236.504	978219.854	978221.206
20	11.500	978253.706	978237.101	978238.460
21	12.000	978271.630	978255.073	978256.438
22	12.500	978290.272	978273.765	978275.135
23	13.000	978309.626	978293.170	978294.549
24	13.500	978329.685	978313.282	978314.658
25	14.000	978350.444	978334.097	978335.464
26	14.500	978371.897	978355.606	978356.972
27	15.000	978394.037	978377.805	978379.177

Table 2. Locally modeled normal gravity field values from the geodetic reference systems

Station No	Longitude (Deg)	Latitude (Deg)	GRS- 30 (mGal)	GRS- 67 (mGal)	WGS- 84 (mGal)
1	8.3215	4.9600	978087.494	978070.446	978072.001
2	8.3410	5.1306	978090.180	978073.140	978074.694
3	8.3215	4.9600	978087.494	978070.446	978072.001
4	8.3410	5.1306	978090.180	978073.140	978074.694
5	8.3330	5.1166	978089.956	978072.915	978074.470
6	8.3483	5.1466	978090.437	978073.397	978074.952
7	8.3410	5.1306	978090.180	978073.140	978074.694
8	8.3446	5.1550	978090.572	978073.532	978075.087
9	8.3446	5.1616	978090.678	978073.639	978075.194
10	8.3600	5.1700	978090.813	978073.775	978075.329
11	8.5160	5.1783	978090.947	978073.909	978075.464
12	8.3216	4.9600	978087.494	978070.446	978072.001
13	8.3410	5.1306	978090.180	978073.140	978074.694
14	8.3416	5.1333	978090.223	978073.183	978074.738
15	8.3366	5.0400	978088.742	978071.698	978073.253
16	8.3140	5.1306	978090.180	978073.140	978074.694
17	8.3366	5.1050	978089.771	978072.730	978074.284
18	8.3433	5.0966	978089.638	978072.596	978074.150
19	8.3450	5.0866	978089.479	978072.437	978073.991
20	8.3466	5.0783	978089.347	978072.305	978073.859
21	8.3498	5.0703	978089.220	978072.177	978073.732
22	8.3528	5.0626	978089.099	978072.055	978073.610
23	8.3548	5.0546	978088.972	978071.929	978073.483
24	8.3566	5.0633	978089.110	978072.067	978073.621
25	8.3666	5.0370	978088.695	978071.651	978073.205
26	8.3450	5.0350	978088.664	978071.619	978073.174
27	8.3466	5.0360	978088.679	978071.635	978073.190
28	8.3366	5.0333	978088.637	978071.593	978073.147
29	8.3140	5.1306	978090.180	978073.140	978074.694
30	8.3140	5.1306	978090.180	978073.140	978074.694
31	8.3588	5.1996	978091.292	978074.255	978075.809
32	8.3613	5.2138	978091.523	978074.486	978076.041
33	8.3596	5.2088	978091.442	978074.405	978075.959
34	8.3610	5.2380	978091.917	978074.882	978076.436
35	8.3595	5.2490	978092.097	978075.062	978076.617
36	8.3576	5.2610	978092.294	978075.259	978076.814
37	8.3593	5.2710	978092.458	978075.424	978076.979
38	8.3583	5.2696	978092.435	978075.401	978076.956
39	8.3610	5.2920	978092.804	978075.771	978077.326
40	8.3598	5.2980	978092.904	978075.871	978077.425
41	8.3561	5.3060	978093.036	978076.003	978077.558
42	8.3526	5.3145	978093.177	978076.144	978077.699
43	8.3363	5.3213	978093.289	978076.257	978077.812
44	8.3403	5.3206	978093.278	978076.246	978077.800
45	8.3336	5.3443	978093.672	978076.641	978078.196

46	8.2750	5.3385	978093.575	978076.544	978078.099
47	8.3200	5.3438	978093.664	978076.633	978078.187
48	8.3150	5.3348	978093.514	978076.482	978078.037
49	8.3410	5.1306	978090.180	978073.140	978074.694
50	8.3410	5.1306	978090.180	978073.140	978074.694
51	8.3453	5.1560	978090.588	978073.549	978075.103
52	8.3483	5.1533	978090.544	978073.505	978075.060
53	8.3300	5.1511	978090.509	978073.470	978075.024
54	8.3266	5.1690	978090.797	978073.759	978075.313
55	8.3075	5.1750	978090.894	978073.856	978075.410
56	8.3086	5.1793	978090.964	978073.925	978075.480
57	8.3005	5.1825	978091.015	978073.977	978075.532
58	8.2950	5.1921	978091.171	978074.133	978075.688
59	8.2870	5.1940	978091.201	978074.164	978075.718
60	8.2781	5.1958	978091.231	978074.193	978075.748
61	8.2701	5.1926	978091.179	978074.141	978075.696
62	8.2601	5.1868	978091.085	978074.047	978075.602
63	8.2528	5.1843	978091.044	978074.006	978075.561
64	8.2455	5.1808	978090.988	978073.950	978075.504
65	8.2361	5.1775	978090.934	978073.896	978075.451
66	8.2193	5.1801	978090.976	978073.938	978075.493
67	8.2198	5.1825	978091.015	978073.977	978075.532
68	8.2115	5.1871	978091.090	978074.052	978075.606
69	8.2000	5.1883	978091.109	978074.071	978075.626
70	8.1955	5.1836	978091.033	978073.995	978075.550
71	8.1905	5.1668	978090.762	978073.723	978075.278
72	8.1831	5.1713	978090.834	978073.796	978075.350
73	8.1758	5.1676	978090.775	978073.736	978075.291
74	8.1653	5.1680	978090.781	978073.742	978075.297
75	8.1570	5.1666	978090.759	978073.720	978075.274
76	8.1481	5.1668	978090.762	978073.723	978075.278
77	8.1838	5.1688	978090.794	978073.755	978075.310
78	8.3410	5.1306	978090.180	978073.140	978074.694
79	8.3410	5.1306	978090.180	978073.140	978074.694
80	8.5283	5.1755	978090.902	978073.864	978075.418
81	8.5540	5.2321	978091.821	978074.785	978076.340
82	8.5675	5.2730	978092.491	978075.457	978077.012
83	8.5425	5.2721	978092.477	978075.442	978076.997
84	8.5270	5.1493	978090.480	978073.441	978074.995
85	8.5273	5.1030	978089.739	978072.698	978074.253
86	8.4931	5.0851	978089.455	978072.413	978073.967
87	8.5230	5.0688	978089.197	978072.154	978073.708
88	8.3410	5.1306	978090.180	978073.140	978074.694
89	8.5000	5.0781	978089.344	978072.301	978073.856
90	8.3386	5.0790	978089.358	978072.316	978073.870
91	8.3806	5.0741	978089.280	978072.238	978073.792
92	8.4798	5.0681	978089.186	978072.143	978073.697

93	8.4636	5.0513	978088.920	978071.877	978073.431
94	8.4408	5.0343	978088.653	978071.608	978073.163
95	8.3410	5.1306	978090.180	978073.140	978074.694
96	8.3410	5.1306	978090.180	978073.140	978074.694
97	8.4453	5.0468	978088.849	978071.805	978073.360
98	8.5008	5.0333	978088.637	978071.593	978073.147
99	8.4880	5.3035	978092.994	978075.962	978077.516
100	8.5405	5.0255	978088.515	978071.470	978073.024
101	8.4513	5.0210	978088.444	978071.399	978072.954
102	8.4501	5.0186	978088.406	978071.361	978072.916
103	8.4666	5.0000	978088.116	978071.070	978072.624
104	8.4666	5.0616	978089.083	978072.040	978073.594
105	8.3488	5.0693	978089.205	978072.162	978073.716
106	8.3378	5.0721	978089.249	978072.206	978073.761
107	8.3238	5.0638	978089.118	978072.074	978073.629
108	8.3070	5.0561	978088.996	978071.953	978073.507
109	8.3488	5.0693	978089.205	978072.162	978073.716
110	8.2500	5.0000	978088.116	978071.070	978072.624
111	8.2250	5.0000	978088.116	978071.070	978072.624
112	8.2000	5.0000	978088.116	978071.070	978072.624
113	8.1500	5.0000	978088.116	978071.070	978072.624
114	8.3410	5.1306	978090.180	978073.140	978074.694
115	8.3410	5.1306	978090.180	978073.140	978074.694
116	8.3606	4.9718	978087.677	978070.630	978072.184
117	8.3655	4.9290	978087.015	978069.967	978071.521
118	8.3776	5.1965	978091.242	978074.204	978075.759
119	8.3938	4.9606	978087.503	978070.456	978072.010
120	8.3883	4.9525	978087.378	978070.330	978071.884
121	8.3953	4.9468	978087.290	978070.242	978071.796
122	8.4051	4.9391	978087.171	978070.122	978071.677
123	8.5683	4.9366	978087.132	978070.084	978071.638
124	8.4223	4.9368	978087.135	978070.087	978071.641
125	8.4215	4.9530	978087.385	978070.338	978071.892
126	8.4386	4.9335	978087.085	978070.036	978071.590
127	8.4473	4.9480	978087.308	978070.260	978071.815
128	8.4551	4.9233	978086.928	978069.879	978071.433
129	8.4611	4.0855	978075.137	978058.057	978059.611
130	8.3410	5.1306	978090.180	978073.140	978074.694
131	8.4641	4.9116	978086.748	978069.699	978071.253
132	8.5230	4.7930	978084.951	978067.897	978069.451
133	8.5188	4.7971	978085.013	978067.958	978069.513
134	8.5120	4.8020	978085.086	978068.032	978069.587
135	5.5096	4.8111	978085.223	978068.169	978069.723
136	8.5105	4.8196	978085.350	978068.297	978069.852
137	8.5075	4.8266	978085.456	978068.403	978069.957
138	8.5046	4.8460	978085.749	978068.697	978070.251
139	8.5026	4.8652	978086.040	978068.988	978070.543

140	8.4921	4.8778	978086.231	978069.181	978070.735
141	8.4775	4.8891	978086.404	978069.353	978070.908
142	8.4696	4.9056	978086.656	978069.606	978071.161
143	8.3410	5.1306	978090.180	978073.140	978074.694
144	8.3410	5.1306	978090.180	978073.140	978074.694
145	8.3613	5.1375	978090.291	978073.251	978074.805
146	8.3566	5.1370	978090.283	978073.243	978074.797
147	8.3640	5.1366	978090.276	978073.236	978074.791
148	8.3670	5.1360	978090.267	978073.227	978074.781
149	8.3410	5.1306	978090.180	978073.140	978074.694
150	8.3410	5.1306	978090.180	978073.140	978074.694
151	8.1780	5.1674	978090.771	978073.733	978075.287
152	8.1806	5.0800	978089.374	978072.332	978073.886
153	8.1848	5.1571	978090.606	978073.566	978075.121
154	8.1894	5.1487	978090.470	978073.431	978074.986
155	8.1943	5.1419	978090.361	978073.321	978074.876
156	8.1922	5.1339	978090.233	978073.193	978074.747
157	8.1892	5.1262	978090.110	978073.069	978074.624
158	8.1877	5.1176	978089.972	978072.931	978074.486
159	8.1826	5.1090	978089.835	978072.794	978074.348
160	8.1845	5.1022	978089.727	978072.685	978074.240
161	8.1903	5.1022	978089.727	978072.685	978074.240
162	8.1964	5.0899	978089.531	978072.489	978074.044
163	8.2033	5.0843	978089.442	978072.400	978073.954
164	8.2110	5.0789	978089.357	978072.314	978073.869
165	8.2183	5.0724	978089.254	978072.211	978073.765
166	8.2206	5.0649	978089.135	978072.092	978073.646
167	8.2246	5.0577	978089.021	978071.978	978073.532
168	8.2180	5.0511	978088.917	978071.873	978073.428
169	8.2327	5.0433	978088.794	978071.750	978073.305
170	8.2360	5.0345	978088.656	978071.611	978073.166
171	8.2430	5.0292	978088.573	978071.528	978073.083
172	8.2503	5.0890	978089.517	978072.475	978074.029
173	8.2568	5.0174	978088.388	978071.343	978072.897
174	8.2617	5.0103	978088.276	978071.231	978072.786
175	8.2533	5.0031	978088.164	978071.118	978072.673
176	8.2735	4.9957	978088.048	978071.002	978072.557
177	8.2764	4.9885	978087.936	978070.890	978072.445
178	8.2789	4.9840	978087.866	978070.820	978072.374
179	8.3410	5.1306	978090.180	978073.140	978074.694
180	8.3410	5.1306	978090.180	978073.140	978074.694
181	8.1743	5.1704	978090.820	978073.781	978075.336
182	8.1745	5.1776	978090.936	978073.898	978075.452
183	8.1751	5.1810	978090.991	978073.953	978075.507
184	8.1713	5.1905	978091.145	978074.107	978075.662
185	8.1614	5.9626	978104.568	978087.566	978089.120
186	8.1542	5.2009	978091.313	978074.276	978075.831

187	8.1471	5.2068	978091.409	978074.372	978075.927
188	8.1404	5.2124	978091.500	978074.463	978076.018
189	8.1283	5.2157	978091.554	978074.517	978076.072
190	8.1165	5.2206	978091.634	978074.597	978076.152
191	8.1030	5.2256	978091.715	978074.679	978076.233
192	8.0899	5.2317	978091.815	978074.779	978076.333
193	8.0841	5.2363	978091.890	978074.854	978076.409
194	8.0667	5.0855	978089.461	978072.419	978073.974
195	8.0570	5.2637	978092.338	978075.304	978076.859
196	8.0388	5.2853	978092.694	978075.660	978077.215
197	8.2820	5.2944	978092.844	978075.811	978077.366
198	8.3410	5.1306	978090.180	978073.140	978074.694
199	8.3410	5.1306	978090.180	978073.140	978074.694
200	8.1773	5.1027	978089.735	978072.693	978074.248
201	8.1724	5.0986	978089.669	978072.628	978074.182
202	8.1673	5.0919	978089.563	978072.521	978074.075
203	8.1613	5.0845	978089.445	978072.403	978073.958
204	8.1561	5.0773	978089.331	978072.289	978073.843
205	8.1536	5.0707	978089.227	978072.184	978073.738
206	8.1572	5.0638	978089.118	978072.074	978073.629
207	8.3410	5.1360	978090.267	978073.227	978074.781
208	8.3410	5.1360	978090.267	978073.227	978074.781
209	8.3319	5.3391	978093.585	978076.554	978078.109
210	8.3408	5.3417	978093.629	978076.598	978078.152
211	8.3473	5.3481	978093.735	978076.705	978078.259
212	8.3380	5.3493	978093.755	978076.725	978078.279
213	8.3605	5.3517	978093.795	978076.765	978078.319
214	8.3664	5.3517	978093.795	978076.765	978078.319
215	8.3731	5.3495	978093.759	978076.728	978078.283
216	8.3808	5.3497	978093.762	978076.731	978078.286
217	8.3886	5.3486	978093.744	978076.713	978078.268
218	8.3945	5.3502	978093.770	978076.740	978078.294
219	8.4018	5.3509	978093.782	978076.751	978078.306
220	8.3855	5.3400	978093.600	978076.569	978078.124
221	8.3782	5.3342	978093.504	978076.472	978078.027
222	8.3827	5.3300	978093.434	978076.402	978077.957
223	8.3902	5.3259	978093.366	978076.334	978077.889
224	8.3972	5.3202	978093.271	978076.239	978077.794
225	8.4012	5.3170	978093.218	978076.186	978077.741
226	8.3943	5.3197	978093.263	978076.231	978077.785
227	8.3900	5.3186	978093.245	978076.213	978077.767
228	8.3842	5.3170	978093.218	978076.186	978077.741
229	8.3794	5.3133	978093.157	978076.124	978077.679
230	8.3761	5.3079	978093.067	978076.035	978077.589
231	8.2217	5.3033	978092.991	978075.958	978077.513
232	8.3640	5.2993	978092.925	978075.892	978077.447
233	8.3612	5.2946	978092.847	978075.814	978077.369

234	8.3410	5.1306	978090.180	978073.140	978074.694
235	8.3410	5.1306	978090.180	978073.140	978074.694
236	8.4691	4.9118	978086.751	978069.702	978071.256
237	8.4752	4.9151	978086.802	978069.752	978071.307
238	8.4796	4.9213	978086.897	978069.848	978071.402
239	8.4801	4.9213	978086.897	978069.848	978071.402
240	8.4890	4.9315	978087.054	978070.005	978071.560
241	8.4971	4.9346	978087.101	978070.053	978071.607
242	8.5042	4.9346	978087.101	978070.053	978071.607
243	8.5118	4.9412	978087.203	978070.155	978071.709
244	8.5206	4.4935	978080.607	978063.542	978065.096
245	8.5294	4.9449	978087.260	978070.212	978071.767
246	8.5382	4.9474	978087.299	978070.251	978071.805
247	8.5401	4.9450	978087.262	978070.214	978071.768
248	8.5454	4.9453	978087.266	978070.218	978071.773
249	8.5510	5.0100	978088.272	978071.226	978072.781
250	8.5570	4.9432	978087.234	978070.186	978071.740
251	8.5631	4.9434	978087.237	978070.189	978071.743
252	8.5694	4.9449	978087.260	978070.212	978071.767
253	8.5755	4.9394	978087.175	978070.127	978071.682
254	8.5827	4.9337	978087.088	978070.039	978071.594
255	8.9264	4.9304	978087.037	978069.988	978071.543
256	8.6000	4.9359	978087.121	978070.073	978071.628
257	8.6091	4.9371	978087.140	978070.092	978071.646
258	8.3410	5.1306	978090.180	978073.140	978074.694

Table 3. Descriptive Statistics of normal gravity field from regional modeling

Variable	N	Mean	SE Mean	StDev	Minimum	Q1	Median	Q3	Maximum
GRS30	27	978184	20.4	106	978055	978088	978162	978272	978394
GRS67	27	978167	20.4	106	978038	978071	978145	978255	978378
WGS84	27	978169	20.4	106	978040	978072	978146	978256	978379

Table 4. Descriptive statistics of normal gravity field from local modeling

Variable	N	Mean	SE Mean	StDev	Minimum	Q1	Median
GRS30mGal	258	978090	0.158	2.53	978075	978089	978090
GRS67mGal	258	978073	0.158	2.54	978058	978072	978073
WGS84mGal	258	978074	0.158	2.54	978060	978073	978075
WGS84-GRS67mGal	258	1.5545	0.0000312	0.000500	1.5540	1.5540	1.5540

Variable	Q3	Maximum	Skewness
GRS30mGal	978091	978105	-0.22
GRS67mGal	978074	978088	-0.22
WGS84mGal	978076	978089	-0.22
WGS84-GRS67mGal	1.5550	1.5550	0.09

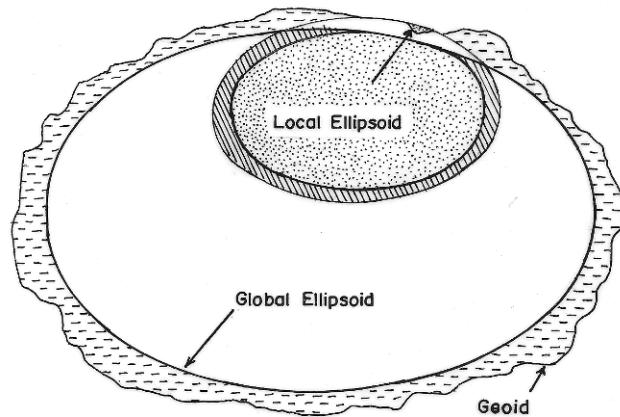


Figure 1a. The earth as an ellipsoid

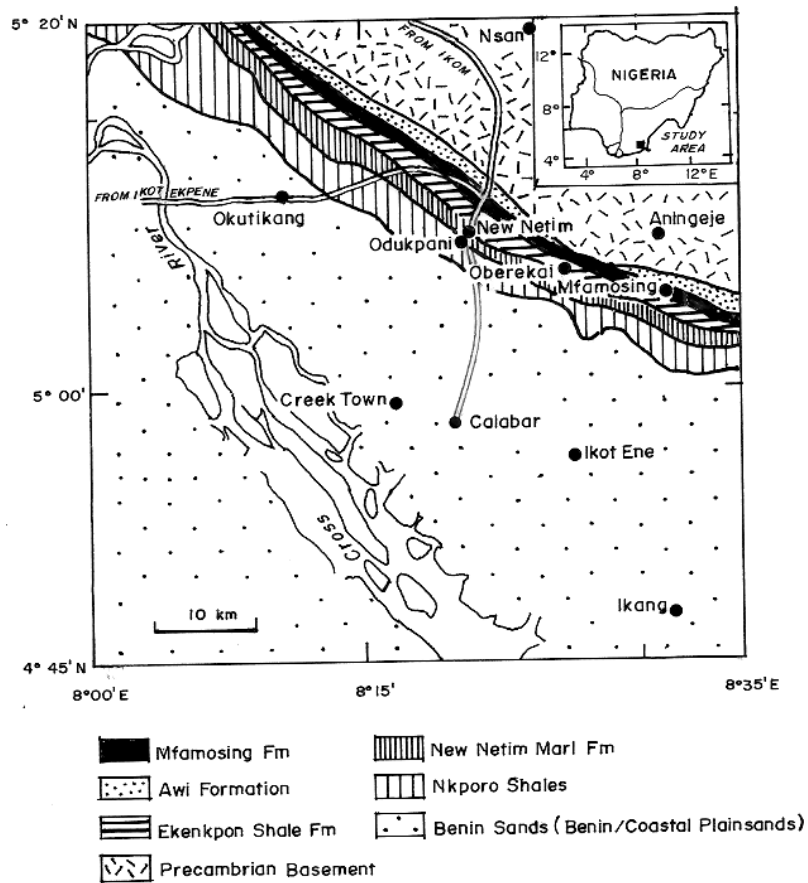


Figure 1b. Geological map of the Calabar Flank, Southeastern Nigeria

