

Mount Kakoulima: An Overview and Analysis

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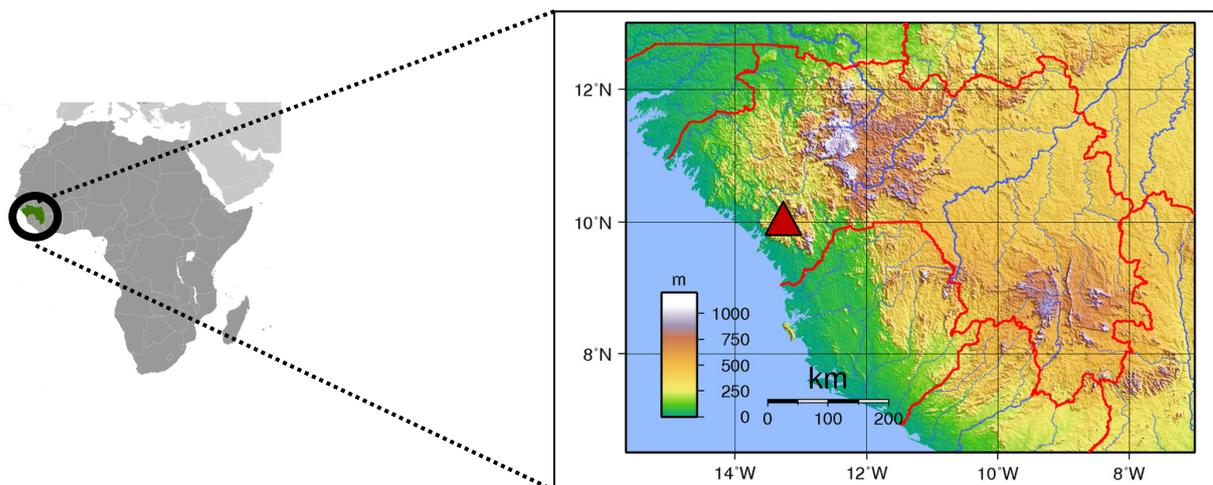
Abstract

The interest generated by mafic-ultramafic rocks is justified by the fact that their structure and formational processes not only have scientific importance but also can lead to metal concentrations, making them economically viable. Located 50 km north-east of Conakry, Guinea's capital, Mount Kakoulima has long been one of the mafic-ultramafic rock areas of the world but has yet to be exploited; thus limited data have been published concerning it. The goal of this investigation is to describe the structure and formation process of Mount Kakoulima by means of geological field observations. The findings indicate that Mount Kakoulima is a stratified solid rock, formed with the result of some magma injections in some intrusive phases. With a Lithological sequence progressing from peridotite (dunite) to pyroxenite to extensive gabbro. The result will improve the understanding of geological processes on Mount Kakoulima to guide future surveys on mineral ores in the scientific and academic community for research and studies in the field of earth sciences.

Keywords: Mount Kakoulima, intrusive complex, mafic-ultramafic, Guinea

1. Introduction

Located 50 km north-east of the capital of Republic of Guinea (Conakry), Kakoulima Complex (Latitude: 9.76694, Longitude: -13.445 and Altitude: 1011 m) dominates the vast plain of lower Guinea. A large part of its slopes is covered by secondary growth vegetations. Frequent fogs, periodic edaphic drought and atmospheric humidity are the essential factors that dominate the ecology of these ridges (Schnell, 1950).



▲ Mount Kakoulima

Figure 1. Map of Africa highlighting Guinea

Source: <http://www.eoearth.org/>

Figure 2. Map of Guinea highlighting the location of Mt Kakoulima

Source: <http://www.eoearth.org/>

The importance in mafic-ultramafic rocks is proven by the fact their study in their various associations leads to two essential problems: the original and current constitution globe, and the genesis and differentiation of magma. In addition, the structure and formational process of mafic-ultramafic rocks may lead to metal concentrations which give them a great economic interest. This was related by several authors in their studies, such as Moores (1973); Gebauer et al. (1981); Wilson (1982); Naldrett (1989); Ferrsrio et al. (1990); Sun et al. (1991); Quadt (1992); Barnes (1994); Gebauer (1996); Garuti et al. (1997); Puchtel et al. (1999); Okano et al. (2000); Dean (2002); Schoenberg et al. (2003); Cawthorn and McKenna (2006); Pirajno et al. (2008); Cawthorn and Lewis (2009).

Virtually most of the world class nickel-Platinum Group Element is produced from mafic-ultramafic complexes. The deposits containing iron, nickel, copper, and Platinum Group Elements (PGE) occur as sulphide concentrations associated with a variety of mafic and ultramafic rocks (Eckstrand et al., 2004; Naldrett, 2004).

In Guinea, these rocks constitute the base of the Peninsula of Kaloum (Kaloum Igneous Complex [KIC]) and are stretched towards the continent where they form the solid mass of Mount Kakoulima (Sylla, 1995). The drilling wells AF0007 to AF0012 which were executed and analyzed by Mega uranium (2004) on the North Grid of the Mount Kakoulima had successfully proven in confirming the original theory that the base of the KIC dips to the south along its northern contact. It dips roughly around 45° inward, and had the capacity to provide traps for concentrating sulphides. The rock types supported in the KIC are consistent with other mineralized mafic-ultramafic intrusions and the contamination of KIC by gneissic wall rocks is a positive feature.

Mount Kakoulima (hereafter referred to as *MK*), one of the world's mafic-ultramafic areas for an extended period of time has yet to be exploited in this context. Chromite and massive sulphides are known in thin subhorizontal horizons near the base of MK (Sable Mining, 2010). Exceptionally, due to favorable natural conditions (high rainfall, high humidity even in the dry season), MK is encircled by dense, picturesque vegetation and controls the rainfall in the Kaloum peninsula.

From the point of view of structure, there are possible analogues of styles among the Great Dyke in Zimbabwe, the Bushveld Igneous Complex in South Africa, and the Stillwater Igneous Complex in Montana (USA). Furthermore, Mega Uranium Ltd. via Marketwire News Releases (2006) confirmed that "there is the potential for massive sulphides similar to the Jinchuan nickel-copper deposit in China".

Due to inadequate resources and support for local research teams, little scientific investigations on MK were realized. Despite these setbacks, Lacroix (1905) described the Central Atlantic Magmatic Province (CAMP) in Guinea, Diallo et al. (1992) discussed the relations with the doleritic sills of Western Guinea and the Central Atlantic rifting, Deckart et al. (2005) studied Geochemistry and Sr, Nd, Pb isotopic composition of the CAMP in Guyana and Guinea, and some Mining Companies recognized iron, nickel, copper, cobalt, chrome and PGE mineralizations on MK.

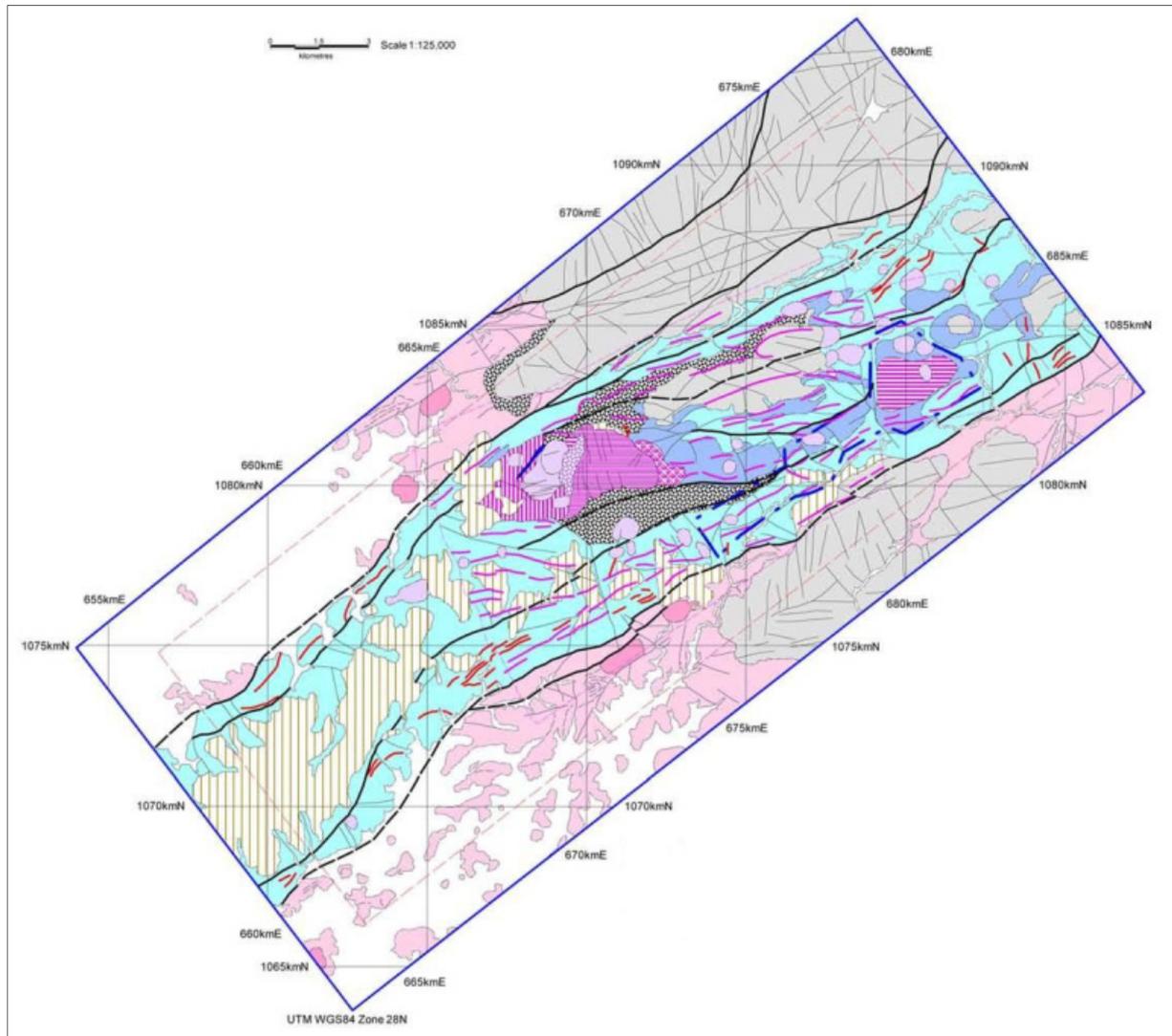
The scope of this paper will help contribute to an understanding of the structure and formational process of MK. The investigative methods involved geological field observations. The findings will improve the understanding of geological processes on the intrusive complex as a guide to further surveys on metals and minerals. Also, this article presents the current state of knowledge on the geological and structural complex intrusive of MK that will be accessible to the scientific and academic community interested in the field of earth sciences. It is hoped that this paper will encourage further studies and research on the issue.

2. Geological Basement

The majority of basement in West Africa was formed during three major orogenic events, these are (as cited by Deckart et al., 2005): Liberian (~2.7 Ga, Hurley et al., 1971), Eburnian (~2.1 Ga=Trans-Amazonian) affecting the Birimian terrains (Abouchami et al., 1990; Ligéois et al., 1991; Boher et al., 1992) and Pan-African (mainly 750-660 and 650-580 Ma; Liégeois et al., 1994). The oldest rocks of Guinea belong to the south-western edge of the West African craton - the Kenema-Man Shield (Archean to Palaeoproterozoic) shared between Guinea, Sierra Leone, Liberia and Ivory Coast, and adjacent Pan-African belts of Rockelides and Bassarides, which are thrust upon the craton and partly covered by the Palaeozoic Bove Basin (Deckart et al., 2005; Villeneuve & Corné, 1994).

The Archean core is characterized by gneisses, granite-gneiss, granites and then by typical association granites-greenstone which are affected by various orogenic of which the last, Liberian orogeny, (2.75 billion years) rejuvenated the existing craton with a very advanced metamorphism having like results, the installation of the large granitoid solid masses (Brinckmann et al., 2008) and offering an environment that may have been suitable for the formation of massive sulphide deposits. In other words, the gites are commonly found in Archean

cratons stable. These Archean cratons are privileged establishment of large stratiform complexes such as MK. The single biotite grains from intrusive formation from the Kakoulima area yielded plateau age of 200.4 ± 0.2 Ma (Deckart et al., 1997; Nomade et al., 2007).



Legend

	Q Surficial deposits		slp Undifferentiated slumped debris and scree
	F Ferricrete		ii Late mafic or ultrabasic intrusion
	scr Scree from Mt. Kakoulima		mch3 Late Mt. Kakoulima magma chamber cumulates
	d Ultrabasic dyke		mch2 Earlier Mt. Kakoulima magma chamber cumulates
	Pz Early Palaeozoic sandstone		mch1 First Mt. Kakoulima magma chamber cumulates
	gr Archean granite		mch? Possible magma chamber cumulates
	mig Archean granite-gneiss, migmatite		ubh Prominent ultrabasic intrusion
	Dyke interpreted from Aster image		ulb Dolerite and ultrabasic intrusions, gneiss screens
	Significant fault		Area of magnetic survey
	Other fault or fracture		Area of concession
	Trace of foliation		

Figure 3. Map of Mt Kakoulima showing geological basement (Source: Sable Mining, 2010)

Geologically, MK belongs to the Kaloum Peninsula and Islands of Loos. The base of the Kaloum Peninsula primarily consists of ultrabasic intrusive rock (dunite). This intrusion in the form of dykes, primary age post (Mesozoic) continued under the Atlantic Ocean, is covered along the coast by marine sedimentary deposits, namely sand, clay and vase, and of altered minerals (produced deterioration-colluviums). The north-eastern end of the solid mass dunite is bordered by the gabbro and norite (Mz) of MK, the granite of the base Precambrian (Pz) and the horizontal sandstones of the Ordovician (Sylla, 1995). MK is a voluminous mafic-ultramafic laccolithic layered intrusion (Diallo et al., 1992). The emplacement of these sills and the laccolith was controlled by old NW-SE lineaments reactivated during Mesozoic time (Bertrand & Villeneuve, 1989; Deckart et al., 2005).

3. Geophysical Surveys

Some geophysical surveys have been done on MK and mineralizations in iron, nickel, copper, cobalt, chrome and PGE have been recognized.

In 1961, a Czechoslovak mission detected concentrations of nickel in its oxidized form near the surface. The study area included three mountains with respective altitudes of 1011 m, 946 m and 878 m. An anomaly was detected by geophysical surveying near Kakoulima train station, which had a depression of 2.6 km². Geophysical survey of the anomaly was not pushed beyond 20 m, although the thickness of the nickeliferous rocks exceeds 24 m. Mineralized surface fields: 2.6 km²; ore tonnage: 38,816,000 tonnes; Nickel content: 0.7-0.9%; Nickel reserves: 185,000 tonnes (Brinckmann et al., 2008). Over the period 1979-1981, the Geosurvey International Society (US) conducted a mineral exploration campaign in the country and the geophysical methods used included airborne and ground magnetics, spectrometry and electromagnetics. Thus, this Geosurvey work had put in evidence that the mafic and ultramafic rocks of MK have one of the most important fields of minerals in Guinea, which supported the enriched minerals such as: copper, nickel, cobalt and PGE (CPG, 1983).

MK was explored by SEMAFO from 1996 to 1998 and by Rio Tinto from 1999 to 2001, with both companies focusing on narrow high grade Ni-Cu-Co-PGE intersections in an upper gabbro/pyroxenite horizon. This resulted in the delineation of an average 0.8 m thick massive sulphide horizon of axial dimensions 100 m x 100 m with weighted average grades of 2.78% Ni and 0.86% Cu. To the south of MK lie extensive laterites overlying mafic-ultramafic cumulates. Nine short vertical holes were drilled in the past by Semafo in a nickeliferous/chromiferous lateritic zone in two areas. All holes mention the omnipresence of Cr with enriched zones of +2% Cr in many instances. One hole intersected a horizon from 11.5m to 26.3m containing 3.28% Cr, 0.58% Ni and 0.13% Co. The Ni/Cr laterite mineralization appears usually to relate to a magnetic anomaly which extends along the southern flank of MK. The zone is a few hundred meters in width (Sable Mining, 2010).

In 2004, Mega Uranium Ltd. drilled a total of 9303.5 meters on 23 diamond drill holes in three phases of exploration, comprising 16 holes testing the northern contact zone and 7 in the southern area. Drilling findings included 14.7 meters 0.35% Ni, 0.39% Cu, 0.17 g/t Pt and 0.79 g/t Pd, and maximum values over one meter intervals of 0.45% Ni and 0.47% Cu (Mega Uranium Ltd., 2006).

In 2010, Sable Mining had a majority interest in a new 298km² areas on MK, which was a prospective for Fe, Ni, Co, Cr, and PGE. Surface sampling and shallow pitting over the general area of the concession had confirmed the geochemistry of the laterite demonstrated in previous investigations. Furthermore, studies also demonstrated why the laterites were exploited for iron in the environs in the 1950s, with many of the results from that program showing +40% of iron (Sable Mining, 2010). Sable Mining goal was to study mineral accumulations in the overlying laterite and saprolite horizons.

Past studies thus far have been mainly focused on the south-east flank of MK intrusive, also concentrating on an area of highly anomalous magnetics extending approximately 13 km in a northeasterly direction. This contains a zone freshly identified by Sable Mining's reinterpretation (both satellite imagery and existing geophysical data) as a further mafic-ultramafic intrusive known as Kakoulima-2. A reconnaissance drilling program with a 160 mm auger was used to drill 50 holes covering over the 20 km² area. A further 17 holes were planned to complete that initial program. All analyses were undertaken by ALS Global's ISO certified laboratory in Brisbane, Australia (Sable Mining, 2011). Initial results from the Kakoulima-2 are tabulated below.

Table 1. Kakoulima-2 results with significant intersection

Hole	Hole depth (m)	Iron Intercept Grade (m)	% Fe ¹	Chromite Intercept Grade (m)	% Cr ₂ O ₃ ²	Nickel Intercept Grade (m)	% Ni ³
Kak11-50	20	20	38.8				
Kak11-48	17			4	2.92		
Kak11-33	16	13	36.7				
Kak11-32	13	13	46.9				
Kak11-31	10	10	43.2	10	4.20		
Kak11-30	29	29	45.1	29	4.14	12	0.63
Kak11-29	17	17	43.0	17	3.93	4	0.57

¹ Generally, samples with Fe content greater than 30% are regarded as superior.

² Generally, samples with Cr₂O₃ content greater than 2.4% are regarded as superior.

³ Generally, samples with Ni content greater than 0.5% are regarded as superior.

Source: Sable Mining, 2011

A mineralogical study by ALS, including QEMSCAN, was initiated to characterize mineralization in different lateritic layers, as a first step to scoping a metallurgical process to produce saleable products. Hand-panning of the lower saprolitic clays, where anomalous magnetic susceptibility was observed, has shown significant quantities of magnetite and chromite in panned concentrates. It was then assumed that the magnetite is a secondary product of the decomposition of the underlying dunite (olivinite) ultramafic rock. In the lateritic sequence above the saprolite, the magnetite altered to hematite, goethite and maghemite (Sable Mining, 2011).

4. Methods

4.1 Literature Review

Data collection took the form of both literature review and field observations. The literature review was primarily aimed at collecting data on geological basement and all geophysical surveys of MK. Documentation included maps, photographs and organization reports. Newspapers also gave insight into the history of activities on the peak. Books, papers and articles mentioning Mt. Kakoulima were also consulted as well as any relevant scholarly works or research dealing with mafic-ultramafic rock formations.

4.2 Materials

During the field observations, a number of important tools were required for data collection:

- An Estwing brand rock hammer was used to gather data on the composition of rock at MK, as well as to determine the nature, strength and history of the rocks. This allowed the researcher to easily split the rocks and expose a fresh surface for analysis.
- The researchers also used maps and photographs to understand the topography and geological basement of MK. (The topographical map was scaled to 1:25,000).
- A geological compass was used together with the above mentioned maps in order to measure the orientation and geometry of the geological structure.
- A loupe with 10x magnification was used to observe details in the rock composition.
- A geological field book was essential for recording observations in the field.

4.3 Field Work

In geosciences, field work is considered fundamental to investigating the development of geological formations over time. As Rudwick (1996) asserts, geology is a science in which fieldwork is a central element of practice, not least because so many important geological features are not mobile. It demands integration and study of content knowledge, observation, theories and analyses as well as experimentation which should form the basis for a coherent and consistent interpretation of the location. Field work was therefore critical to this research as observations of the structure, composition and nature of MK, provided the primary data needed to explain its geological formation and lithology.

Observations of the intrusive mafic-ultramafic formations on MK were taken by a team of researchers during several excursions between March and May, 2010.

5. Results and Discussion

5.1 Structure of Mount Kakoulima

Being the nearest large mountain to the capital Conakry, the MK is one of the most exceptional rock formations in West Africa. Field research proved that on the crest of MK, is a remarkable image on the rock appearance; an enormous outline of a dog’s head which the locals dub as smoking dog (in French ‘*le chien qui fume*’) although it is not as impressive as it seems to be. The MK represents a stratified laccolite in Figure 4 (also confirmed by Diallo et al., 1992), which intersects older granitic and metamorphic rocks, covered by Paleozoic sedimentary deposits. At the base of MK, granite and dunite appear to be in direct contact.

From a geomorphological perspective, the slopes of MK form a terrace as shown in Figure 5 below. This terrace reflects the layering structure of the internal mass (stratigraphy), which is related to the different rock types that are encountered in horizontal layers.

From the base to top, field survey exposed the following formation: peridotite (dunite), pyroxenite and gabbro-norite, this is illustrated below in Figure 6. Also, geophysical surveys have proven that MK is stratified mafic and ultramafic rocks, encountered by sulphide mineralization within pyroxenite and gabbro horizon. The average thicknesses of rocks are: peridotite (dunite): 100 m; pyroxenite: 150 m; and gabbro-norite upper 250 m.



Figure 4. Showing a laccolite form
Source: <http://www.mmah.e-monsite.com>



Figure 5. Showing slopes forming a terrace
Source: <http://www.mmah.e-monsite.com>

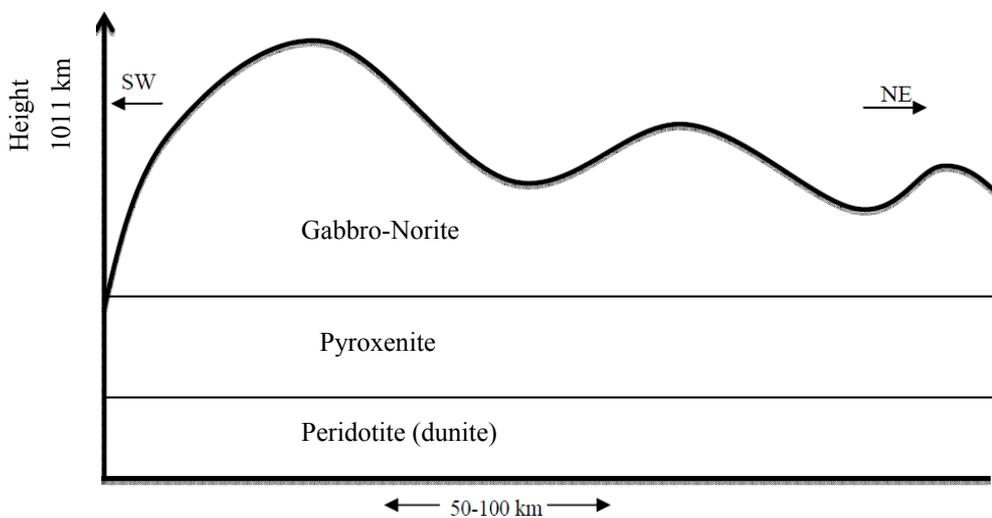


Figure 6. Illustration of the form and structure of Mt Kakoulima

MK from a structural point of view, is possibly analogue when comparing researches done on the Great Dyke in Zimbabwe (Wilson, 1982; Wilson, 1992; Wilson, 1996; Wilson et al., 2000), the Bushveld Igneous Complex in South Africa (Von et al., 1985; Mitchell, 1990; Kruger, 1990; Ealesaand & Cawthornb, 1996), the Stillwater Igneous Complex in Montana (Bow et al., 1982; Raedeke & McCallum, 1984; McCallum, 1996), and the Munni Munni Complex in Australia (Barnes & Hoatson, 1994). Each of these intrusions is a kind of assemblage of magmatic sediments forming stratified layers where sediment is represented by mineral grains from a gravitational fall in the magma. Cumulus minerals (Plagioclase, orthopyroxene, clinopyroxene, olivine) form the backbone of these rocks and the enveloping cement. In other words, these rock cumuli are accumulations of crystals deposited after a more or less prolonged gravitational fall.

The cumulus phase is dominated by the processes of crystal fractionation. The accumulation of minerals decommitted as a result of the crystallization process occurring in the magma chamber due to a high density of crystals. The accumulated rocks generally exhibit well developed textural stratification. After the cumulus phase, subsequent formation of rocks can happen by crystallization of new mineral fusion (intercumulus) and compaction of the cumulus pile. This phase, which includes overgrowth, develops strong compositional zoning. Crystallization partially forms overgrowth on the primary cumulus crystals, and discrete intercumulus minerals. Compaction occurs in response to the increasing pressure acting on the individual crystals. This leads to a rotation of cumulus crystals, local pressure dissolution (resorption), and a filter pressing of the intercumulus liquid.

These phenomena are called crystalline assemblage of structured mineral stages of MK, which are characterized by a bottom to top succession of ultramafic to mafic rocks: peridotite (dunite), pyroxenite, and gabbro-norite.

5.2 Process of Formation of Solid

Field research proved that the dunite (almost olives) lies at the base of the massif. It is an ultramafic rock containing less than 45% SiO₂ formed by early crystallizing and accumulation of grains of olivine and pyroxene by injection of magma during the 1st intrusive phase (see Figure 7). It is also the source of chromite, iron, nickel and cobalt, which is also confirmed by geophysical surveys (see Table 1). On the outcrops of dunite, one can observe large crystals of diallage with the naked eye. The quantity grows in an identical way from bottom to top on the entire thickness of the rock.

The relationships of the dunite with pyroxenite are very complex. These relationships are clearly visible in outcrops under Creek Bridge at the 46.2 km mark along the former Conakry-Kankan railway. Here, dunite with average grains is intersected by a superimposed pyroxenite dyke with fine grains. A little higher along the bed of the same creek, dunite is substituted by this same pyroxenite. Pyroxenite is also an ultramafic rock, composed basically of minerals of the pyroxene. It is formed by accumulation and compaction of pyroxene crystals by injection of magma during the 2nd intrusive phase (see Figure 7). It is also source of chromite cumulates.

In the Kitima river bed, one can observe a pyroxenite injection in the form of dykes in dunite with bitownite. Thus, these observations demonstrate the existence of a relationship break between the first layer of dunite and pyroxenite of the second layer. On the Kitima river bank, not far from the place where dunite is intersected by pyroxenite, the gabbro-norite injections in the form of dykes in gabbro-norite melanocrate are observable. Gabbro-norite is a mafic rock (45-52% SiO₂) which is often associated with iron, titanium and platinum mineralizations. It is formed as cumulates mafic intrusion by injection of magma during the 3rd intrusive phase (see Figure 7).

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